

# **Iodine-129 in Soils from Northern Ukraine and the Retrospective Dosimetry of the Iodine-131 Exposure after the Chernobyl Accident**

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## ABSTRACT

Forty-eight soil profiles down to a depth of 40 cm were taken in Russia and Ukraine in 1995 and 1997, respectively, in order to investigate the feasibility of retrospective dosimetry of the  $^{131}\text{I}$  exposure after the Chernobyl accident via the long-lived  $^{129}\text{I}$ . The sampling sites covered areas almost not affected by fallout from the Chernobyl accident such as Moscow/Russia and the Zhitomir district in Ukraine as well as the highly contaminated Korosten and Narodici districts in Ukraine.  $^{129}\text{I}$  was analyzed by radiochemical neutron activation analysis (RNAA) and accelerator mass spectrometry (AMS).  $^{127}\text{I}$  was measured for some profiles by RNAA or ion chromatography. The results for  $^{127}\text{I}$  demonstrated large differences in the capabilities of the soils to store iodine over long time-spans. The depth profiles of  $^{129}\text{I}$  and of  $^{137}\text{Cs}$  showed large differences in the migration behavior between the two nuclides but also for each nuclide among the different sampling sites. Though it cannot be quantified how much  $^{129}\text{I}$  and  $^{137}\text{Cs}$  was lost out of the soil columns into deeper depths, the inventories in the columns were taken as proxies for the total inventories. For  $^{129}\text{I}$ , these inventories were at least three orders of magnitude higher than a pre-nuclear value of  $0.084 \pm 0.017 \text{ mBq m}^{-2}$  derived from a soil profile taken in 1939 in Lutovinovo/Russia. From the samples from Moscow and Zhitomir a pre-Chernobyl  $^{129}\text{I}$  inventory of  $(44 \pm 24) \text{ mBq m}^{-2}$  was determined, limiting the feasibility of  $^{129}\text{I}$  retrospective dosimetry to areas where the  $^{129}\text{I}$  inventories exceed  $100 \text{ mBq m}^{-2}$ . Higher average  $^{129}\text{I}$  inventories in the Korosten and Narodici districts of  $130 \text{ mBq m}^{-2}$  and  $848 \text{ mBq m}^{-2}$ , respectively, allowed determination of the  $^{129}\text{I}$  fallout due to the Chernobyl accident. Based on the total  $^{129}\text{I}$  inventories and on literature data for the atomic ratio of  $^{129}\text{I}/^{131}\text{I} = 13.6 \pm 2.8$  for the Chernobyl emissions and on aggregated dose coefficients for  $^{131}\text{I}$ , the thyroid exposure due to  $^{131}\text{I}$  after the Chernobyl accident was estimated for the inhabitants of 4 villages in the Korosten and of 3 villages in the Narodici districts. The limitations and uncertainties of the  $^{129}\text{I}$  retrospective dosimetry are discussed.

## 1 Introduction

The long-lived radionuclide  $^{129}\text{I}$  ( $T_{1/2} = 15.7 \text{ Ma}$ ,  $E_{\beta, \text{max}} = 0.2 \text{ MeV}$ ,  $E_{\gamma} = 39.58 \text{ keV}$ ,  $I_{\gamma} = 0.0752$ ) is produced in nature by cosmic-ray induced spallation of xenon in the atmosphere and by spontaneous fission in the geosphere. The total natural inventory of  $^{129}\text{I}$  in the lithosphere was estimated to be 50 000 kg (327 TBq) and of this only 263 kg (1.7 TBq) was estimated to be available as “free” inventory of atmosphere, hydrosphere and biosphere and mixing with the stable  $^{127}\text{I}$ ; see Schmidt et al. (1998) for details. The natural abundances of  $^{129}\text{I}$  have been globally and in a sustainable way changed by  $^{129}\text{I}$  released by man into the environment. Man-made  $^{129}\text{I}$  originates primarily from fission of  $^{235}\text{U}$  and  $^{239}\text{Pu}$  with isobaric yields for fission induced by thermal neutrons of 0.68 % and 1.6 %, respectively; iodine is produced in thermal fission of  $^{235}\text{U}$  with isotopic ratios of  $^{131}\text{I}/^{129}\text{I} = 3.82$  and  $^{129}\text{I}/^{127}\text{I} = 6.06$  (England and Rider 1994).

Estimates of the  $^{129}\text{I}$  releases from explosions of nuclear weapons are between 43 kg (0.28 TBq) and 150 kg (0.98 TBq) (Carter and Moghissi, 1977; UNSCEAR 1982; Chamberlain, 1991; Eisenbud and Gesell, 1997). Schmidt et al. (1998) estimated that the Chernobyl accident released less than 2 kg (0.013 TBq) of  $^{129}\text{I}$ . The vast amount of man-made  $^{129}\text{I}$  in the environment is due to releases from reprocessing plants. For example, about 3500 kg (23 TBq) of  $^{129}\text{I}$  were released from European reprocessing plants up to the year 2000, of this 70 % being by the La Hague reprocessing plant (Raisbeck et al., 1995; Gray et al., 1995; Parker, 2001; Webmaster Cogema, 2001; BNFL, 1999; Groupe radioécologique Nord-Cotentin, 1999). Of the  $^{129}\text{I}$  emissions from La Hague 97 % were liquid discharges, compared with 85 % from Sellafield. Large amounts of man-made  $^{129}\text{I}$  are still stored in spent nuclear fuel, estimated to be 5660 kg (37 TBq) in 1990 (Finkel and Suter, 1993). There is little information about the releases from U.S. military reprocessing plants (Moran et al., 1999); nothing is published about plants in the former USSR and other eastern countries.

Man-made  $^{129}\text{I}$  enters the environmental iodine cycles and changes the natural iodine isotopic abundances. The natural  $^{129}\text{I}/^{127}\text{I}$  ratios in the oceanic mixing layer have been globally changed by two orders of magnitude from about  $1.5 \cdot 10^{-12}$  (Schink et al., 1995; Moran et al., 1998) to about  $10^{-10}$  in areas remote from emissions and with orders of magnitude higher ratios close to the release points (Yiou et al., 1994). Though these changes are of radiological relevance only in the closest proximity to reprocessing plants, they should be carefully monitored and the potential of  $^{129}\text{I}$  as a man-made tracer for environmental processes should be exploited. This requires a detailed understanding of the pre-nuclear baseline values, of the anthropogenic changes, of the radioecology of  $^{129}\text{I}$  and of the environmental chemistry of iodine. However, the radioecology of  $^{129}\text{I}$  is still insufficiently known due to analytical and methodological problems (Schmidt et al., 1998). Systematic investigations to establish reliable analytical protocols for  $^{129}\text{I}$  and  $^{127}\text{I}$  analyses in various environmental materials have since been established to close some gaps in our knowledge about the radioecology of  $^{129}\text{I}$  (Schmidt, 1998; Schmidt et al., 1998; Szidat et al., 2000c; Szidat, 2000; Szidat et al., 2000b; Ernst et al., 2002; Ernst et al., 2003).

The analysis of  $^{129}\text{I}$  in environmental samples, in particular in soils, offers an opportunity for the retrospective dosimetry of the radiation exposure caused by short-lived iodine isotopes, such as  $^{131}\text{I}$ , long after the radionuclides were released in accidents or incidents. There are several requirements for this retrospective dosimetry to be feasible. The  $^{129}\text{I}$  from the accident must not have disappeared from the analyzed soil horizons by migration into deeper soil layers and it must be distinguishable from potentially existing fallout prior to accident. Finally, the radioecological modeling connecting the deposition density of  $^{131}\text{I}$  derived from the observed  $^{129}\text{I}$  deposition densities with the thyroid dose must be reliable.

The mathematical model underlying the retrospective dosimetry of  $^{131}\text{I}$  exposures using  $^{129}\text{I}$  is given by equation 1.

$$H_{thy} = \left( D(^{129}\text{I}) - D_{\text{pre-Ch.}}(^{129}\text{I}) \right) \cdot \frac{A_{131}}{A_{129}} \cdot DC_{131} \cdot f(t) \quad \text{Eq. 1}$$

where

- $H_{thy}$  is the committed equivalent dose to the thyroid due to  $^{131}\text{I}$  in Sv,
- $D(^{129}\text{I})$  the integral  $^{129}\text{I}$  deposition density at a location in  $\text{Bq m}^{-2}$ ,
- $D_{\text{pre-Ch.}}(^{129}\text{I})$  the integral  $^{129}\text{I}$  deposition density from pre-Chernobyl fallout in  $\text{Bq m}^{-2}$ ,
- $\frac{A_{131}}{A_{129}}$  the activity ratio of  $^{131}\text{I}$  and  $^{129}\text{I}$  at the time of the accident,
- $DC_{131}$  an aggregated dose coefficient for  $^{131}\text{I}$  exposure of the thyroid for a single fallout event in  $\text{Sv Bq}^{-1} \text{m}^2$  derived from a radioecological model,
- $f(t)$  a correction factor which accounts for the decay of  $^{131}\text{I}$  between the accident and arrival of the  $^{131}\text{I}$  at the location of the exposure.

In principle,  $^{129}\text{I}$  in equation 1 could be substituted by other radionuclides, such as  $^{137}\text{Cs}$  or  $^{90}\text{Sr}$ , if the emission and atmospheric transport are homogeneous as for the global weapons fallout. In the case of the Chernobyl accident, however, the different elements were chemically differentiated during the explosion and the subsequent emissions and were further differentiated by a time dependent atmospheric transport in the troposphere as plumes with differing elemental compositions; see UNSCEAR (2000) and references therein.

This work is part of an ongoing radioecological project by which the environmental behaviour of various radionuclides released by the Chernobyl accident and the resulting radiation exposure of the inhabitants in the highly contaminated areas in Northern Ukraine are studied (Schmidt, 1998;

Schmidt et al., 1998; Filss et al., 1998; Slavov et al., 1999; Botsch, 2000; Botsch et al., 2000; Handl et al., 2003; Mewis, 2004). Here, we deal with the analysis of  $^{127}\text{I}$  and  $^{129}\text{I}$  in soils from Russia and Northern Ukraine and investigate the feasibility of using  $^{129}\text{I}$  as an indicator for  $^{131}\text{I}$  fallout after the Chernobyl accident. We discuss the various input quantities on which retrospective dose estimates depend and present the results of a retrospective dosimetry of the exposure to  $^{131}\text{I}$  for people living in some villages in the Korosten and Narodici districts of Northern Ukraine. Some preliminary data for  $^{129}\text{I}$  from this work have been reported by Schmidt et al. (1998) and Michel et al. (2000, 2002a, 2002b, 2002c).

There are several earlier publications dealing with the retrospective dosimetry of  $^{131}\text{I}$  exposures using  $^{129}\text{I}$  (Straume et al., 1996; Robl et al., 1997; Mironov et al., 1999, 2002; Reithmeier et al., 2002; Pietrzak-Flis et al., 2003). Most of the earlier work aimed at evaluating the feasibility of  $^{129}\text{I}$  retrospective dosimetry and on determining some of the input quantities of equation 1. Only Robl et al. (1997) and Pietrzak-Flis et al. (2003) did complete evaluations ending up with thyroid dose estimates.

## 2 Experimental

We investigated 2 soil profiles from Moscow/Russia and 42 soil profiles from Northern Ukraine taken in 1995 and 1997, respectively. The sampling sites covered areas almost not affected by fallout from the Chernobyl accident such as Moscow/Russia and Zhitomir/Ukraine as well as moderately and highly contaminated regions of Northern Ukraine in the Korosten and the Narodici districts, respectively (Fig. 1). According to an early categorization (Law of Ukraine, 1991) the sampling sites in Korosten belong to the contamination zone III with  $^{137}\text{Cs}$  deposition densities between  $185 \text{ kBq m}^{-2}$  and  $555 \text{ kBq m}^{-2}$  ( $5 - 15 \text{ Ci km}^{-2}$ ), those in Narodici to the contamination zone II with

$^{137}\text{Cs}$  deposition densities between  $555 \text{ kBq m}^{-2}$  and  $1480 \text{ kBq m}^{-2}$  ( $15 - 40 \text{ Ci km}^{-2}$ ). Zone I is the exclusion zone inside the 30 km cycle around the Chernobyl nuclear power plant (CNPP).

The Zhitomir district is located 160 km to the southwest from the CNPP; this district is considered to be the most weakly contaminated part of the Northern Ukraine. The sampling area included the following settlements: Barashevka, Davidovka, Levkov, and Oserjanka. The Korosten district is situated 120 km west of CNPP which is expected to be a moderately contaminated area of zone III. The samples were collected in the vicinity of the villages Nemirovka, Woronewo, Tschigiri, and Kupetsch. The Narodici district is located at the distance of about 70 km from CNPP in the extreme west of zone II. It represents one of the most highly contaminated areas of Ukraine. The soil profiles were taken in environments surrounding three settlements: Christinovka, Novo Scharno, and Nosdrischtsche.

All sampling sites were chosen at state and private farms at locations without bushes and trees with reportedly constant land use and not affected by human activities during recent decades and which should represent undisturbed soils. At each plot, up to 7 soil profiles randomly distributed within an area of about  $4 \text{ km}^2$  were sampled.

To take the soil profiles  $20 \times 20$ -cm soil blocks were collected. To this end, 45 cm-deep holes with surface areas of  $1 \times 1 \text{ m}$  were dug leaving  $20 \times 20$ -cm soil columns in a corner. These columns were sectioned at 1-cm intervals within a top 5-cm layer, then 5-cm intervals down to 25 cm and a final interval to a depth of 40 cm. After determining the fresh weight, the material was dried at room temperature, milled, and thereafter sieved through a 2-mm mesh sieve; see Handl et al. (2003) for details of sampling.

The samples were analysed for  $^{129}\text{I}$  by radiochemical neutron activation analysis (RNAA) and accelerator mass spectrometry (AMS), for  $^{127}\text{I}$  by RNAA and ion chromatography (IC), and for  $^{137}\text{Cs}$  and  $^{134}\text{Cs}$  by  $\gamma$ -spectrometry.

RNAA turned out to be only capable of determining  $^{129}\text{I}$  in top-soils; see Schmidt et al. (1998) for details. For deeper layers,  $^{129}\text{I}$  could only be determined via AMS. With maximum sample masses of 80 g a detection limit according to ISO 11929-7 (ISO, 2003) of  $1 \mu\text{Bq kg}^{-1}$  and  $1.8 \text{ mBq kg}^{-1}$  for  $^{129}\text{I}$  was achieved by AMS and RNAA, respectively. Typical detection limits of  $^{127}\text{I}$  were  $90 \text{ ng g}^{-1}$  and  $60 \text{ ng g}^{-1}$  for RNAA and IC, respectively (Schmidt et al., 1998). By combining results from AMS and IC analyses  $^{129}\text{I}/^{127}\text{I}$ -ratios in soils were detectable down to  $5 \cdot 10^{-12}$ .

The  $^{129}\text{I}$  blank values of the total analyses were determined with Woodward iodine as trace catcher. Typical blank  $^{129}\text{I}/^{127}\text{I}$  ratios of the total analyses performed in two different laboratories were in terms of geometric means and standard deviations<sup>1</sup>  $(1.3 \times 2.0^{\pm 1}) \cdot 10^{-12}$  and  $(3.5 \times 1.7^{\pm 1}) \cdot 10^{-13}$ , respectively (Ernst et al., 2003).

The experimental relative standard uncertainties of the concentrations of the iodine isotopes were between 2.8 % and 21 % for  $^{127}\text{I}$  and between 4.4 % and 56 % (RNAA), respectively 19 % (AMS) for  $^{129}\text{I}$ . The average relative uncertainties were 9.0 % in the case of  $^{127}\text{I}$  and 9.3 % in the case of  $^{129}\text{I}$ .

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<sup>1</sup> In this paper, we use the convenient notation of e.g.  $1.3 \times 2.0^{\pm 1}$  to present geometric means and geometric standard deviations in analogy to the frequently used notation of e.g.  $1.5 \pm 0.3$  when reporting arithmetic means and standard deviations. For an assumed logarithmic normal distribution, one obtains an interval containing about 66 % of the data by multiplying, respectively by dividing, the geometric mean by the geometric standard deviation.



Detailed descriptions of the sample preparations, the analytical techniques (AMS, RNAA, and IC), and quality control measures and on participation in round robin exercises (Roberts et al., 1997; Roberts and Caffee, 2000) are given elsewhere (Schmidt et al., 1998; Szidat et al., 2000a, 2000b, 2000c; Ernst et al., 2003).

As it was not feasible to analyze all depth profiles in detail for  $^{127}\text{I}$  and  $^{129}\text{I}$ , the deposition densities of 30 depth profiles were determined by analysis of bulked or mixed samples. These were prepared from the individual profile samples so that they allowed calculation of the total activities in each profile. Since for  $^{137}\text{Cs}$  all samples from each depth profile had been measured by gamma-spectrometry (Botsch, 2000; Botsch et al., 2000; Handl et al., 2003) the quality of this procedure could be checked. The ratio of the  $^{137}\text{Cs}$  deposition densities determined from individual soil profile samples to those determined from mixed samples was  $0.99 \pm 0.07$ , thus demonstrating the feasibility of the procedure.

### 3 Experimental results

The numerical results of all the analyses reported in this work can be obtained from the internet at [www.zsr.uni-hannover.de](http://www.zsr.uni-hannover.de). Detailed results for  $^{137}\text{Cs}$  and  $^{134}\text{Cs}$  in the soil profiles dealt with in this paper and for the radiation exposure due to  $^{137}\text{Cs}$  were recently published (Handl et al., 2003).

Prior to this study, a screening investigation had been performed on  $^{129}\text{I}$  and  $^{127}\text{I}$  in 51 samples of top soils ( $d \leq 10$  cm) from various affected and unaffected locations in Ukraine, Belarus and Russia sampled between September 1991 and October 1992 using RNAA (table 1). The distances between the sampling locations of these top soils and the Chernobyl reactor ranged from 0.8 km to more than 400 km. The variability of the  $^{129}\text{I}$  specific activities was much larger than that of the  $^{127}\text{I}$  concentrations and the maximum of the  $^{129}\text{I}$  specific activities exceeded that of measurements of soils from

Northern Germany by an order of magnitude. The results showed clearly that the  $^{129}\text{I}$  fallout from the Chernobyl accident could be identified in the soils. It was also seen that the variability of the  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios in the top soils was much larger than that of the  $^{129}\text{I}$  specific activities. The latter observation could only be explained by differences in the  $^{127}\text{I}$  concentrations of the soils which must not be neglected when interpreting  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios. Therefore, the  $^{127}\text{I}$  concentrations shall be presented first.

### 3.1 Depth profiles of $^{127}\text{I}$

For all those samples measured by RNAA  $^{127}\text{I}$  data are available, i.e. data exist typically for the upper 10 cm. For those samples investigated by AMS, complete analyses of  $^{127}\text{I}$  by IC were only performed for a subset of four depth profiles. The mixed samples were analyzed for  $^{129}\text{I}$  only.

The  $^{127}\text{I}$  concentrations range from  $0.42 \text{ mg kg}^{-1}$  to  $8.0 \text{ mg kg}^{-1}$ , well in agreement with the range observed in the earlier screening measurements of topsoil samples. In Fig. 2, the four complete  $^{127}\text{I}$  concentration depth profiles are shown. The concentrations are fairly uniform with depth for the individual profiles, but show remarkable differences between the four locations. The data vary by about a factor of ten. The variability of the experimental data is not due the experimental standard uncertainties, which ranged from 3 % to 21 % with an average of 9 %.

In the following, we frequently make use of the inventory of  $^{127}\text{I}$  and  $^{129}\text{I}$  within individual depth increments, in units of  $\text{g m}^{-2}$  and  $\text{mBq m}^{-2}$ , respectively, calculated according to equation 2:

$$A_{\text{F}}(d_{\text{max}}, d_{\text{min}}) = \int_{d_{\text{max}}}^{d_{\text{min}}} C(d) \cdot r(d) dd \quad \text{Eq. 2}$$

with  $d_{\max} = 40$  cm being the maximum sampling depth,  $d_{\min}$  the depth of the top of the actual sampling interval,  $C(d)$  the specific activity or mass concentration at depth  $d$  of  $^{129}\text{I}$  and  $^{127}\text{I}$ , respectively, and  $\mathbf{r}(d)$  the bulk density of the air-dried soil at depth  $d$ .

Provided that all fallout nuclides are contained in the sampled depth  $d_{\max} = 40$  cm the total inventory  $D$  is given by

$$D \approx A_{\text{F}}(d_{\max}, 0) = \int_{d_{\max}}^0 C(d) \cdot \mathbf{r}(d) dd \quad . \quad \text{Eq. 3}$$

The depth dependence of the inventories within individual depth increments of  $^{127}\text{I}$  is shown in Fig. 2 for the four soil profiles from Moscow, Nemirovka, and Nosdrischtsche. Interpreting the  $A_{\text{F},127}(d_{\max}, 0)$  as the time integrated total inventory  $D_{127}$  of  $^{127}\text{I}$ , the data from Fig. 2 would imply that there were differences in the  $^{127}\text{I}$  atmospheric input of nearly a factor of ten. This appears rather unlikely. Different iodine concentrations in air and precipitation are only to be expected on large scales over landmasses. Thus, the observed differences can only be attributed to the different soil characteristics and hydrological conditions. Low  $^{127}\text{I}$  concentrations and consequently low total inventories point to a significant loss of stable iodine over the long times of atmospheric input. High  $^{127}\text{I}$  concentrations and high total inventories point to lower, but still unaccountable loss of iodine and therefore should be a better proxy of the real integral deposition density of stable iodine. High total inventories point to a smaller migration into depth and to a more efficient sorption or accumulation in the particular soil, low ones to a higher migration and less efficient sorption or accumulation. We shall come back to this when discussing the  $^{129}\text{I}$  depth profiles.

As a consequence of the different transport processes involved in the migration of iodine in the soils, the long-term fallout of  $^{127}\text{I}$  is expected to result in an exponential decrease of inventories within individual depth increments over the depth of the soil. The  $^{127}\text{I}$  inventories within individual depth increments are coarsely in agreement with this expectation (Fig. 2). They decrease exponen-

tially with characteristic lengths between 19 cm and 24 cm. The total inventories of  $^{127}\text{I}$  range from  $0.39 \text{ g m}^{-2}$  to  $3.4 \text{ g m}^{-2}$  with a geometric mean of  $0.9 \text{ g m}^{-2}$  and a geometric standard deviation of 2.6. The inventories within individual depth increments decrease by factors of about three over the entire depth profiles.

The differences between the total inventories at the different locations indicate that substantial amounts of stable iodine which was input over thousands of years have been lost at the various locations into the deeper, water-saturated soil zones due to individual soil characteristics, hydrological conditions, and hydrodynamic dispersion. With this conclusion it becomes a crucial question whether the water-unsaturated soil zones are sufficient archives to preserve  $^{129}\text{I}$  for the purpose of a retrospective dosimetry of  $^{131}\text{I}$ . This will be discussed below.

### 3.2 Depth profiles of $^{129}\text{I}$

The  $^{129}\text{I}$  specific activities measured in the depth profiles span nearly four orders of magnitude, ranging from  $0.006 \text{ mBq kg}^{-1}$  to  $24.5 \text{ mBq kg}^{-1}$ . Only in two samples of the initial screening investigation of top soils were higher  $^{129}\text{I}$  specific activities observed:  $52 \text{ mBq kg}^{-1}$  in a sample from the proximity of Gomel in Belarus and  $111 \text{ mBq kg}^{-1}$  in a top soil sampled 0.8 km from the Chernobyl reactor.

Examples of the depth dependencies of  $^{129}\text{I}$  specific activities in the differently contaminated areas are shown in Fig. 3. For the areas not significantly affected by Chernobyl fallout they decrease from  $\sim 1 \text{ mBq kg}^{-1}$  at the surface to  $\sim 0.01 \text{ mBq kg}^{-1}$  in the lowest layer.  $^{129}\text{I}$  specific activities decrease nearly monotonically in the two profiles from Moscow and in Levkov 2, while in the profile Levkov 1 a local maximum appears in the layer from 5 to 10 cm. In the profiles from zone III,  $^{129}\text{I}$  specific activities decrease also nearly monotonically: the profile from Woronewo 4 from 10

mBq kg<sup>-1</sup> to 0.1 mBq kg<sup>-1</sup>, those from Nemirovka and Kupetsch-Tschernjanka from values larger than 1 mBq kg<sup>-1</sup> to 0.01 mBq kg<sup>-1</sup>.

The profiles from Nosdrischtsche and Christinovka in zone II look different than the other ones. The profile from Christinovka river shore shows a local maximum of 24.5 mBq kg<sup>-1</sup> in the 2 to 3 cm layer, followed by a monotonous decrease down to 0.1 mBq kg<sup>-1</sup>. The Christinovka meadow profile generally shows relatively low (1.1 mBq kg<sup>-1</sup> to 0.2 mBq kg<sup>-1</sup>) specific activities. After a short decrease from surface, they exhibit a broad maximum extending from 10 cm to 30 cm depth. All the specific activities change by just one order of magnitude over the entire profile while for all the other profiles of all locations the specific activities drop by two orders of magnitude.

### 3.3 <sup>129</sup>I/<sup>127</sup>I isotopic ratios

The <sup>129</sup>I/<sup>127</sup>I isotopic ratios in the four completely analyzed soil profiles are presented in Fig. 4. For Moscow and Nemirovka 2, they decrease monotonically from the surface into the depth. The surface ratios are between  $2 \times 10^{-8}$  and  $5 \times 10^{-8}$ . At depth, they differ by about an order of magnitude between  $3 \times 10^{-10}$  and  $3.5 \times 10^{-9}$ . For the profile Nosdrischtsche 2, the <sup>129</sup>I/<sup>127</sup>I isotopic ratios increase from  $3 \times 10^{-7}$  at the surface to more than  $10^{-6}$  between 10 and 20 cm. Below 20 cm they decrease by more than an order of magnitude to  $2 \times 10^{-8}$ . For all the profiles, even the deepest samples are affected by man-made <sup>129</sup>I. The lowest isotopic ratios at depth exceeded the pre-nuclear marine equilibrium ratios by more than 2 orders of magnitude and the lowest ratio measured so far, a pre-nuclear soil at  $(5.7 \pm 1.1) \cdot 10^{-12}$  (Szidat et al., 2000b) by a factor of 50. The range of <sup>129</sup>I/<sup>127</sup>I isotopic ratios observed in the screening measurements (table 1) is slightly larger than that measured in the four profiles reflecting the wider sampling area.

### 3.4 Integral deposition densities of $^{129}\text{I}$ in Ukraine

Inventories within individual depth increments of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  were calculated from the  $^{129}\text{I}$  and  $^{137}\text{Cs}$  specific activities in the depth profiles according to equation 2 as proxy for the integral deposition densities (table 2). The profiles of these inventories are presented in Fig. 5.

In contrast to the respective data for  $^{127}\text{I}$  (Fig. 2), the inventories within individual depth increments of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  decrease with depth by about two orders of magnitude, much steeper than those of  $^{127}\text{I}$ . Even the profiles from the areas not affected by the Chernobyl accident, i.e. Moscow, Levkov and Barashevka, show this behaviour. This supports the assumption that most of the fallout is still confined to the topmost 40 cm. In the lowest layers (25 to 40 cm) of the depth profiles, we observe only between 1 % and 10 % of the total  $^{129}\text{I}$  and even less in case of  $^{137}\text{Cs}$  (down to 0.2 %).

In table 3 the results of the total inventories of  $^{129}\text{I}$  are summarized and compared with data for a pre-nuclear soil from Lutinovo in Russia (Szidat et al., 2000b) and from a recent investigation of 7 deep depth profiles from Lower Saxony, Germany (Ernst et al., 2003).

All the total inventories are at least three orders of magnitude higher than a value of  $0.084 \pm 0.017$   $\text{mBq m}^{-2}$  derived from a pre-nuclear soil profile taken in 1939 in Lutovinovo/ Russia (Szidat et al., 2000b). The geometric means and standard deviations of the total  $^{129}\text{I}$  inventories are  $49 \times 1.5^{\pm 1}$   $\text{mBq m}^{-2}$  for samples from Moscow,  $38 \times 1.7^{\pm 1}$   $\text{mBq m}^{-2}$  for samples from Zhitomir/Ukraine,  $130 \times 1.5^{\pm 1}$   $\text{mBq m}^{-2}$  for samples from contamination zone III and  $848 \times 1.5^{\pm 1}$   $\text{mBq m}^{-2}$  for those from contamination zone II.

The logarithmic means of the total  $^{137}\text{Cs}$  inventories of  $4.1 \times 1.2^{\pm 1}$   $\text{kBq m}^{-2}$  for Moscow,  $4.6 \times 1.2^{\pm 1}$   $\text{kBq m}^{-2}$  for Zhitomir, and  $4.8 \times 1.4^{\pm 1}$   $\text{kBq m}^{-2}$  for Lower Saxony are well in agreement and confirm

the negligible influence of Chernobyl fallout at these locations. They are dominated by global fallout of  $^{137}\text{Cs}$  from atmospheric nuclear explosions.

For  $^{129}\text{I}$ , the results for the soil profiles from Moscow and Zhitomir are significantly lower than those obtained from seven soil profiles taken in 1999 near Hanover in Lower Saxony/Germany which showed total inventories of  $168 \times 1.1^{\pm 1} \text{ mBq m}^{-2}$  (Ernst et al., 2003). Thus, it appears that Western Europe is more affected by recent fallout of  $^{129}\text{I}$  as a consequence of emissions from the European reprocessing plants in La Hague and Sellafield than Ukraine and Russia. While for Western Europe the ongoing fallout of  $^{129}\text{I}$  has been investigated (Szidat et al., 2000b; Schnabel et al., 2001), for Ukraine or Russia no such investigations exist up to now.

The ratios  $D(^{129}\text{I})/D(^{137}\text{Cs})$  of the total inventories of  $^{129}\text{I}$  and of  $^{137}\text{Cs}$  give some evidence for the origin of environmental  $^{129}\text{I}$ , since  $^{129}\text{I}$  and  $^{137}\text{Cs}$  from nuclear explosions should represent the isobaric fission yields while from reprocessing plants  $^{129}\text{I}$  is preferentially emitted. In Moscow and around Zhitomir, the ratios  $D(^{129}\text{I})/D(^{137}\text{Cs})$  show a geometric mean of  $8.5 \times 10^{-6}$  with a geometric standard deviation of 1.6. In zones III and II, the ratios  $D(^{129}\text{I})/D(^{137}\text{Cs})$  are nearly two orders of magnitude lower. In zone III, they show a geometric mean of  $2.0 \times 10^{-7}$  with a geometric standard deviation of 1.9, and in zone II a geometric mean of  $2.8 \times 10^{-7}$  with a geometric standard deviation of 1.3. The latter two mean ratios are close to what one expects according to equation 4 for the activity ratio  $A_{\text{fission}}(^{129}\text{I})/A_{\text{fission}}(^{137}\text{Cs})$  in fallout from atmospheric weapon tests and in undifferentiated Chernobyl debris.

$$\frac{A_{\text{fission}}(^{129}\text{I})}{A_{\text{fission}}(^{137}\text{Cs})} = \frac{Y(129) \cdot I(^{129}\text{I})}{Y(137) \cdot I(^{137}\text{Cs})} = \frac{0.0065 \cdot 1,4 \cdot 10^{-15} \text{ s}^{-1}}{0.0626 \cdot 7,3 \cdot 10^{-10} \text{ s}^{-1}} = 2,0 \times 10^{-7} \quad \text{Eq. 4}$$

with

$Y(A)$  being the fission yield of the isobars  $A = 129$  and  $A = 137$  and

$\lambda$  the decay constants of  $^{129}\text{I}$  and  $^{137}\text{Cs}$ .

In zone II, the ratios  $D(^{129}\text{I})/D(^{137}\text{Cs})$  are just slightly higher than the fission value and show a small variability. In zone III, they match the fission value on the average, but the individual ratios are spread over an order of magnitude. The nearly two orders of magnitude higher ratios in Moscow and around Zhitomir clearly indicate the existence of another additional source of  $^{129}\text{I}$ , most likely the effect of the European reprocessing plants. The total inventories of  $^{129}\text{I}$  support this interpretation since they are much higher than the total  $^{129}\text{I}$  inventories due to fallout of atmospheric nuclear explosions in the northern hemisphere. Estimate of the weapons fallout of  $^{129}\text{I}$  are between 1.4 mBq  $\text{m}^{-2}$  (Rao and Fehn, 1999) and 2.9 mBq  $\text{m}^{-2}$  (Oktay et al., 2000; Schink et al., 1995).

## 4 Discussion

The discussion will first focus on the migration behavior of the iodine isotopes in soils. Then, we shall discuss the individual input data for a retrospective dosimetry of  $^{131}\text{I}$  exposures according to equation 1.

### 4.1 Migration of Iodine and deposition densities of $^{127}\text{I}$ and $^{129}\text{I}$

As in most migration studies, the soil profiles investigated in this work sample only the upper water-unsaturated soil zone. There, the  $^{127}\text{I}$  concentrations are fairly constant in the individual depth profiles but vary considerably from one location to another. This is in accordance with the results of an investigation of seven deeper soil profiles from Lower Saxony, Germany with depths down to 2.5 m (Ernst et al., 2003). It was also observed that the  $^{129}\text{I}$  concentrations were fairly uniform with depth for the individual profiles in the upper, unsaturated soil zones; they steeply decreased at deeper layers from about 1 mg  $\text{kg}^{-1}$  down to 0.06 mg  $\text{kg}^{-1}$ . This observation was independent of land



use and, since these depth profiles were completely analyzed with respect to their soil characteristics, cannot be explained by different biological activities or bioturbation.

The differences observed in the  $^{127}\text{I}$  concentrations and total inventories can only be explained by differences in the migration behavior in the soil. As discussed by Ernst et al. (2003) in detail, the natural stable iodine concentrations in the soils cannot be explained by the iodine concentrations in the bedrock, but are the result of wet and dry deposition of atmospheric iodine originating from the sea. They are the result of an input over thousands of years since the annual deposition density rates from precipitation are only in the range from  $1 \text{ mg m}^{-2} \text{ a}^{-1}$  to  $6 \text{ mg m}^{-2} \text{ a}^{-1}$  (Szidat et al., 2000b). Over long time spans, the differences in the apparent integral deposition densities of  $^{127}\text{I}$  cannot be explained by variations due to wet and dry deposition as long as one does not compare forested and non-forested landscapes.

The iodine concentrations in the water-unsaturated soil zone are the result of the competition between downward migration and accumulation via upward directed evapo-transpiration. Migration is the movement of iodine from the surface into the deeper layers of the soil as a consequence of various transporting processes. The processes resulting in the migration may also contribute to an accumulation of the iodine in certain upper layers of the soil. Accumulation of a substance, contained in precipitation and consequently transported into the soil water as dissolved material, as particulate matter or as colloids, is an increase of its concentration in the water-unsaturated zone (upper soil) by evaporation of water. The evaporation process forces the water to move into the opposite direction of the migration, without the latter being excluded. Diffusion boundary layers, which are formed from adsorbed water and capillary water at the soil matrix, contribute to the accumulation and different modes of physical and chemical sorption which occur on widely differing time scales influence the accumulation and migration. Iodine which reaches the water-saturated soil zones disappears from the system via ground leachate and water transport and, consequently, the differing

$^{127}\text{I}$  integral deposition densities can well be explained by loss of the stable iodine from the soil over long time spans.

The question whether  $^{129}\text{I}$  from recent fallout was lost by migration from the sampled depth interval is essential for the retrospective dosimetry. From the existing data of  $^{129}\text{I}$  we cannot obtain evidence of what losses might occur on short time scales of tens of years. However, the data of Fig. 5 indicate that the pre- and post-Chernobyl man-made  $^{129}\text{I}$  is still confined to the topmost 40 cm, but we cannot quantify how much of the  $^{129}\text{I}$  has been lost. The investigation of the soil profiles from Lower Saxony, Germany demonstrated that man-made  $^{129}\text{I}$  has reached a depth of 2.5 m (Ernst et al., 2003). The lowest observed  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios were  $\sim 10^{-10}$ . However, between 77 % and 96 % of the  $^{129}\text{I}$  in the soil profiles from Lower Saxony was found in the topmost 40 cm. For  $^{127}\text{I}$ , only between 17 % and 81 % were in the upper 40 cm.

A comparison of the depth dependencies of the inventories within individual depth increments of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  demonstrates significant differences between the migration behaviours of the two radionuclides. This is shown for the profiles from areas not affected by the Chernobyl fallout in Fig. 5. The depth dependencies of  $^{137}\text{Cs}$  generally are steeper than those of  $^{129}\text{I}$  pointing to a higher mobility of the latter nuclide. The two profiles from Levkov are quite similar in the case of  $^{137}\text{Cs}$ , but differ considerably for  $^{129}\text{I}$ . Also for the upper part of the soil profile from Barashevka different functional behaviors are seen. For the profiles from Moscow and for Barashevka 1 no  $^{137}\text{Cs}$  could be detected in the 25 cm to 40 cm layer.

Also the  $^{129}\text{I}$  depth profiles differ considerably for soils from zones III and II (Fig. 5). For the lowest soil layer sampled, the  $^{129}\text{I}$  inventories within individual depth increments show a much larger spread for samples from zone III than the total inventories, thus emphasizing the variability in migration at the different locations. Extremes are the depth profiles from Christinovka meadow and

river shore. However, they also show that the migration of  $^{137}\text{Cs}$  and  $^{129}\text{I}$  run roughly parallel (Fig. 5). However, the  $^{129}\text{I}$  profiles have smaller gradients with depth than those of  $^{137}\text{Cs}$ . At the same time, Fig. 5 shows that the completely different migration situation at Christinovka meadow also affects  $^{137}\text{Cs}$  in the same way, i.e. that the migration is faster in this location. The soil profile for Christinovka meadow is exceptional. Deviations of the  $^{137}\text{Cs}/^{134}\text{Cs}$  activity ratio from the ratio of  $1.88 \pm 0.05$  at the time of the Chernobyl accident (Handl et al., 2003) give evidence that even caesium from the Chernobyl accident has already reached the 25 cm to 40 cm depth region in the profile Christinovka meadow and that the  $^{137}\text{Cs}$  from nuclear weapons fallout has already disappeared from the sampled soil column. Also a loss of  $^{129}\text{I}$  appears likely for this soil profile.

The  $^{129}\text{I}/^{127}\text{I}$  depth dependencies of the different profiles (Fig. 4) are more difficult to interpret because they reflect a mixture of long-term effects for  $^{127}\text{I}$  and of short-term ones for  $^{129}\text{I}$ . However, they can be understood by comparing them with the  $^{127}\text{I}$  inventories within individual depth increments (compare to Fig. 2). The observations in this work are consistent with those made by Ernst et al. (2003). A slower decrease with depth of the  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios is observed at locations with lower total  $^{127}\text{I}$  inventories and faster iodine migration. We observe the steepest  $^{129}\text{I}/^{127}\text{I}$  decrease by two orders of magnitude in the profile from Nemirovka II which had the highest total  $^{127}\text{I}$  inventory, i.e. in the soil which accumulated the iodine best over even long time scales. For soil Moscow VII, which had a comparably low total  $^{127}\text{I}$  inventory, the  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios decrease just by one order of magnitude. The soil Nosdrischtsche 2 had the smallest total  $^{127}\text{I}$  inventory pointing to a high migration and a low accumulation of iodine. Well in line with this, its  $^{129}\text{I}/^{127}\text{I}$  ratios exhibit their maximum at depth, the Chernobyl  $^{129}\text{I}$  already being substantially dislocated to depth.

In summary, one may conclude that a sampling depth of 40 cm appears to be sufficient to cover most of the recent fallout of  $^{129}\text{I}$  and that the inventories in the sampled soil columns reveal the  $^{129}\text{I}$  deposition densities. But, in view of the results from Lower Saxony, Germany (Ernst et al., 2003),

there remains a caveat. These data can only be regarded as lower limits of the total inventories. In spite of that, we shall use them as a proxy for the total inventories.

## 4.2 Pre-Chernobyl inventory of $^{129}\text{I}$

The feasibility of a retrospective  $^{131}\text{I}$  dosimetry via  $^{129}\text{I}$  also depends crucially on the question of whether the  $^{129}\text{I}$  from the Chernobyl fallout can be detected against the background of the pre-Chernobyl  $^{129}\text{I}$  deposition as the result of the global weapon testing fallout and of liquid and aerial  $^{129}\text{I}$  discharges from reprocessing plants and subsequent atmospheric transport. It was realized that this is a real problem when comparing the  $^{129}\text{I}$  specific activities with results for soils from Lower Saxony, Germany, where 25 samples of topsoil showed logarithmic means of  $^{129}\text{I}$  specific activities and of  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios of  $1.7 \text{ mBq kg}^{-1}$  and  $5.9 \cdot 10^{-8}$  with geometric standard deviations of 1.7 and 3.6, respectively.

In this work, we can estimate the pre-Chernobyl total inventory from the two soil profiles from Moscow, Russia and from 12 soil profiles from the vicinity of Zhitomir, Ukraine. Based on the total  $^{137}\text{Cs}$  inventories and missing  $^{134}\text{Cs}$  in these profiles, both regions are practically unaffected by the Chernobyl fallout and thus can serve for such an estimate. From the Moscow soils we estimate  $D_{\text{pre-Ch.}}(^{129}\text{I}) = (51 \pm 19) \text{ mBq m}^{-2}$  (geometric mean and standard deviation  $49 \text{ mBq m}^{-2} \times 1.5^{\pm 1}$ ), while the soils from Zhitomir give  $D_{129,\text{pre-Ch.}} = (47 \pm 27) \text{ mBq m}^{-2}$  (geometric mean and standard deviation:  $38 \text{ mBq m}^{-2}$  and 1.7). Combining our data from Moscow and from Zhitomir district we take  $D_{\text{pre-Ch.}}(^{129}\text{I}) = (44 \pm 24) \text{ mBq m}^{-2}$  as our best estimate for the pre-Chernobyl total inventory of  $^{129}\text{I}$ .

This value of  $44 \text{ mBq m}^{-2}$  is much higher than that estimated for the global weapons fallout inventory of  $^{129}\text{I}$  for which the estimates lie between  $1.4 \text{ mBq m}^{-2}$  and  $2.9 \text{ mBq m}^{-2}$  (Schink et al., 1995;

Rao and Fehn, 1999; Oktay et al., 2000). Thus, it turns out that also in Eastern Europe a substantial amount of  $^{129}\text{I}$  from emissions from reprocessing plants has been deposited.

For Eastern Europe just a few estimates of the pre-Chernobyl inventories exist (Straume et al., 1996; Reithmeier et al., 2002; Mironov et al., 1999, 2002). Robl et al. (1997) and Pietrzak-Flis et al. (2003) assumed the same pre-Chernobyl contribution to  $^{129}\text{I}$  as Straume et al. (1996).

Straume et al. (1996) investigated the feasibility of  $^{131}\text{I}$  retrospective dosimetry via  $^{129}\text{I}$ . They analyzed soil samples from Belarus to a depth of 30 cm using AMS. They estimated the pre-Chernobyl  $^{129}\text{I}$  in soils by analyzing a sample from San Joaquin valley in California, USA which showed a  $^{129}\text{I}$  atom concentration of  $6 \times 10^7 \text{ g}^{-1}$ . Assuming this value as typical for a depth of 30 cm and a soil with a density of  $1.5 \text{ g cm}^{-3}$  this atom concentration is equivalent to a total inventory of  $26 \text{ mBq m}^{-2}$ . The relative good agreement with our estimate appears to us accidental, since neither the latitude dependence of the atmospheric weapons fallout was considered nor the even more important influences of emissions from reprocessing plants which are different in Europe and in North America.

Reithmeier et al. (2002) reported some data on  $^{129}\text{I}$  concentrations and  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios for six lakes (Ammersee, Lago Maggiore, Lago di Garda, Baikal, Balkhash, and Issyk-kul) and discussed  $^{131}\text{I}$  dose reconstruction via  $^{129}\text{I}$  in the former Soviet Union. They claimed that their data allowed estimation of pre-Chernobyl inventory of  $^{129}\text{I}$  of  $(7.4 \pm 1.3) \times 10^{12} \text{ atoms m}^{-2} \cong (10.4 \pm 1.6) \text{ mBq m}^{-2}$  (Lake Baikal) and  $(8.5 \pm 3.1) \times 10^{12} \text{ atoms m}^{-2} \cong (11.9 \pm 4.3) \text{ mBq m}^{-2}$  (Lake Issyk-kul). Their estimates are significantly lower than our data for the pre-Chernobyl inventories in Russia and the Northern Ukraine.

Mironov et al. (1999, 2002) used  $^{129}\text{I}$  and  $^{137}\text{Cs}$  to estimate the total  $^{131}\text{I}$  inventories due to Chernobyl in Belarus. From the analysis of a pre-Chernobyl soil sample taken 400 km north of Chernobyl

in 1985 and assuming a maximum penetration depth of 20 cm, both for  $^{129}\text{I}$  and  $^{137}\text{Cs}$ , they derived pre-Chernobyl inventory of  $26 \text{ mBq m}^{-2}$  and  $2.4 \text{ kBq m}^{-2}$  for  $^{129}\text{I}$  and  $^{137}\text{Cs}$ , respectively. They noted that the  $^{129}\text{I}/^{137}\text{Cs}$  ratio can only be explained if there is an additional source of  $^{129}\text{I}$  in Belarus except the atmospheric weapons fallout. Their value is well within the range of data observed in our work.

The pre-Chernobyl inventory measured or estimated for Eastern Europe are different from those measured in Western Europe at locations far away from the reprocessing plants. Lower Saxony just had minor fallout from the Chernobyl accident and thus should also be suitable to estimate the pre-Chernobyl fallout of  $^{129}\text{I}$ . However, we found for Lower Saxony  $^{129}\text{I}$  deposition densities with a geometric mean and standard deviation of  $168 \text{ mBq m}^{-2} \times 1.5^{\pm 1}$  (Ernst et al., 2003), much higher than in Moscow or Zhitomir. This is due to the fact that in Lower Saxony there is a continuous and ongoing fallout of  $^{129}\text{I}$  due to the releases from Sellafield and La Hague reaching values up to  $\sim 20 \text{ mBq m}^{-2} \text{ a}^{-1}$  via precipitation at the end of the 1990s (Szidat et al., 2000b; Michel et al., 2002b, 2003).  $^{129}\text{I}$  fallout increased from  $0.014 \text{ mBq m}^{-2} \text{ a}^{-1}$  in 1950 to  $\sim 0.8 \text{ mBq m}^{-2} \text{ a}^{-1}$  in 1984/1985 in Switzerland (Wagner et al., 1996). From 1997 – 2000 the annual deposition of  $^{129}\text{I}$  in Lower Saxony ranged between 2 and  $19 \text{ mBq m}^{-2} \text{ a}^{-1}$ . There are no data available, which allow judging how important this fallout is for Poland, Belarus, Ukraine, and Russia. Also, there is no evidence to judge the importance of releases from Russian reprocessing plants. Thus, at the moment we are urged to use the estimates from Moscow and Zhitomir.

The disadvantage of the earlier estimates is that they do not account for the uncertainty of the pre-Chernobyl inventory. The uncertainty  $u(D_{\text{pre-Ch.}}(^{129}\text{I}))$  finally decides about the capability to distinguish between pre-Chernobyl fallout and that due to the accident. The decision threshold  $D_{129}^*$  which decides whether a measured inventory  $D(^{129}\text{I})$  exceeds that of the pre-Chernobyl fallout is according to ISO 11929-7 (ISO, 2003) given by

$$D_{129}^* = k_{1-\alpha} \cdot \sqrt{2} \cdot u(D_{\text{pre-Ch.}}(^{129}\text{I})) \quad \text{Eq. 5}$$

$u(D_{\text{pre-Ch.}}(^{129}\text{I}))$  is the standard uncertainty of  $D_{\text{pre-Ch.}}(^{129}\text{I})$  according to the ISO Guide for the Expression of Uncertainty in Measurements (ISO, 1995). Thus, the question of whether a retrospective dosimetry is feasible depends on how accurately the pre-Chernobyl fallout can be determined.

For the probability of the error of 1<sup>st</sup> kind  $\alpha = 0.05$  one obtains  $k_{1-\alpha} = 1.65$  and with  $u(D_{\text{pre-Ch.}}(^{129}\text{I})) = 24 \text{ mBq m}^{-2}$  a decision threshold  $D_{129}^* = 56 \text{ mBq m}^{-2}$  for the net inventory of  $^{129}\text{I}$  due to the Chernobyl fallout ( $D(^{129}\text{I}) - D_{\text{pre-Ch.}}(^{129}\text{I})$ ). Only if measured  $^{129}\text{I}$  inventories exceed  $100 \text{ mBq m}^{-2}$  ( $=D_{\text{pre-Ch.}}(^{129}\text{I}) + D_{129}^*$ ) one cannot conclude that they are due to fallout from the Chernobyl accident. This leads us to the conclusion that only in contamination zones III and II  $^{129}\text{I}$  from the Chernobyl accident can be seen (Fig. 6). Moreover, the data obtained for zones III and II have to be corrected for the pre-Chernobyl fallout of  $44 \text{ mBq m}^{-2}$ .

### 4.3 Inventories of $^{129}\text{I}$ due to the Chernobyl accident

The measured total inventories of  $^{129}\text{I}$  are plotted against those of  $^{137}\text{Cs}$  in Fig. 6. For comparison, we included the data reported by Mironov et al. (2002) and from Straume et al. (1996) for soils from Belarus. In addition, the logarithmic means and standard deviations of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  total inventories from 7 depth profiles from Lower Saxony/Germany (Ernst et al., 2003) are shown.

The data by Straume et al. (1996) and by Mironov et al. (2002) exhibit the same trends as ours. Moreover, they fill a gap in our data set for total  $^{137}\text{Cs}$  inventories between  $10 \text{ kBq m}^{-2}$  and  $200 \text{ kBq m}^{-2}$  where a number of samples appears to be significantly affected by the  $^{129}\text{I}$  fallout from the Chernobyl accident. The large variation of the  $^{129}\text{I}$  data together with the likewise extreme variability of the total  $^{129}\text{I}$  inventories for  $^{137}\text{Cs}$  integral deposition densities above  $1 \text{ MBq m}^{-2}$  observed by

them and Straume et al. (1996) demonstrates clearly that a retrospective dosimetry of the  $^{131}\text{I}$  exposure via  $^{137}\text{Cs}$  is not meaningful.

This is well in line with the actual differences between total  $^{131}\text{I}$  and the  $^{137}\text{Cs}$  inventories which occurred due to time-dependent elemental fractionation during emission and atmospheric transport. UNSCEAR (2000) reported values between 10 and 30 for this ratio and Talerko (2004) calculated ratios of  $^{131}\text{I}$  and  $^{137}\text{I}$  inventories between 5 and 28 for Ukraine taking into account the detailed emission history and the actual weather conditions.

Also Hou et al. (2003) investigated  $^{129}\text{I}$  and  $^{137}\text{Cs}$  in 11 soils from Belarus and Russia. Their data are fairly consistent with ours for  $^{137}\text{Cs}$  inventories exceeding  $100 \text{ kBq m}^{-2}$ . But, there is a problem for estimating the pre-Chernobyl inventory: their lowest  $^{129}\text{I}$  inventory is about  $10 \text{ mBq m}^{-2}$  for one sample with a  $^{137}\text{Cs}$  inventory of only  $0.91 \text{ kBq m}^{-2}$ . The latter value is unreasonably low since the fallout of the nuclear weapons tests is about  $5 \text{ kBq m}^{-2}$ .

Fig. 6 emphasizes the geographical variability of the  $D_{\text{pre-Ch.}}(^{129}\text{I})$  if one compares the data from Lower Saxony with those measured in soils from Russia and Ukraine. Consequently, it is a *conditio sine qua non* for the retrospective dosimetry via  $^{129}\text{I}$  to derive reliable data and an uncertainty estimate for  $D_{\text{pre-Ch.}}(^{129}\text{I})$ . From Fig. 6, one also observes that the  $D_{\text{pre-Ch.}}(^{137}\text{Cs})$  of  $(4.8 \times 1.4^{\pm 1}) \text{ kBq m}^{-2}$  derived for Lower Saxony/Germany (Ernst et al., 2003) is also a good estimate for the pre-Chernobyl integral  $^{137}\text{Cs}$  inventory in Ukraine, Belarus and Russia as it is to be expected from the homogeneity of the global weapons fallout.

In contamination zones II and III, the  $^{129}\text{I}$  inventories are significantly higher than those from Moscow and Zhitomir and allow estimation of the contribution from the Chernobyl fallout by correcting the radionuclide data for the pre-Chernobyl fallout. There remains the general tendency of the  $^{129}\text{I}$



inventories to increase with the  $^{137}\text{Cs}$  inventories, but the large variability of the  $^{129}\text{I}$  inventory for a given one of  $^{137}\text{Cs}$  prohibits substituting  $^{129}\text{I}$  by  $^{137}\text{Cs}$  for the purpose of retrospective dosimetry. This is due to the differentiation between  $^{137}\text{Cs}$  and  $^{129}\text{I}$  during the atmospheric dispersion and the fact that the districts of Korosten und Narodici received the fallout not from the first plume which went north affecting in particular Belarus, where Straume et al. (1996) observed the up to now highest  $^{129}\text{I}$  inventories.

Also, the individual villages in zones II and III investigated in this work show considerable variations of the total inventories both for  $^{129}\text{I}$  (table 3) and  $^{137}\text{Cs}$  (Handl et al., 2003). In order to account for the large variability of the total  $^{129}\text{I}$  inventories, we have calculated the expectation values of the total inventories for each village assuming a logarithmic normal distribution of the total inventories. With an arithmetic mean  $\mathbf{m}$  and a standard deviation  $\mathbf{s}$  of the natural logarithms of the total inventories, the expectation value  $E(D(^{129}\text{I}))$  is given by

$$E(D(^{129}\text{I})) = \exp(\mathbf{m} + \mathbf{s}^2/2) \quad . \quad \text{Eq. 6}$$

With the expectation values of the total  $^{129}\text{I}$  inventories those of  $^{131}\text{I}$  were calculated in order to derive dose estimates for the people living in these villages. However, the  $^{131}\text{I}/^{129}\text{I}$  activity ratio at the time of the accident has to be discussed first.

#### 4.4 The $^{131}\text{I}/^{129}\text{I}$ activity ratio

There are several estimates of the  $^{129}\text{I}/^{131}\text{I}$  isotopic ratio at the time of the accident (table 4). Straume et al. (1996) estimated the  $^{129}\text{I}/^{131}\text{I}$  isotopic ratio on the basis of  $^{129}\text{I}$  measurements of four aliquots of one soil sample from Gomel for which an  $^{131}\text{I}$  measurements was available and obtained  $^{129}\text{I}/^{131}\text{I} = 12 \pm 3$ ; individual values ranged from 10 - 14. This isotopic ratio is equivalent to an activity ratio

$^{131}\text{I}/^{129}\text{I}$  of  $5.9 \times 10^7$ . Mironov et al. (2002) used the same method and obtained  $^{129}\text{I}/^{131}\text{I} = 15.2 \pm 4.7$  from the analysis of 24 soil samples. VanMiddlesworth and Handl (1997) investigated animal thyroid glands and obtained  $^{129}\text{I}/^{131}\text{I} = 27 \pm 10$ , the individual values ranging from 18.6 to 89.3. Szidat et al. (2000c) analyzed sewage sludge and found  $^{129}\text{I}/^{131}\text{I} = 34 \pm 2.8$ . From the analysis of rain samples from Israel and from Munich, Germany a weighted mean of  $^{129}\text{I}/^{131}\text{I} = 19 \pm 5$  was obtained, the individual samples ranging from 9 to 35 (Paul et al., 1987; Kutschera et al., 1988). Also the results of burn-up calculations are conflicting. Kirchner and Noack (1988) estimated  $^{129}\text{I}/^{131}\text{I} = 22.8$ , while Ermilov et al. (1993) estimated a range between 11 and 15 for  $^{129}\text{I}/^{131}\text{I}$ .

There are different sources of uncertainties in the various estimates. One is related to the operational history of the reactor and to the accuracy achievable by burn-up calculations. A second one is that the amount of pre-Chernobyl  $^{129}\text{I}$  which might be in an environmental sample is not known. A third one is that large corrections for decay of  $^{131}\text{I}$  during the long time-spans of the emission and post-emission phases affect the accuracy of the ratio determinations; see Talerko (2004) for a detailed discussion.

The first source of uncertainties affects the results of Ermilov (1993) and of Kirchner and Noack (1988) which are highly discrepant. The rain samples (Paul et al., 1987; Kutschera et al., 1988), the sewage sludge (Szidat et al., 2000c), and the animal thyroids VanMiddlesworth and Handl (1996) are likely to be affected by uncertainty sources numbers two and three. The  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios in bovine thyroid glands from Austria and in deer thyroid glands from Bavaria/Germany sampled in May/June 1986 clearly represent a peak due to the Chernobyl accident over the long-term trend of iodine isotopic ratios in Western Europe (Michel et al., 2000). Given, however, the uncertainties of the arrival of the radioactive cloud, of the efficiency of the transfer of fallout iodine into the thyroids, and of the potential addition to the Chernobyl  $^{129}\text{I}$  from the accident of other anthropogenic  $^{129}\text{I}$ , the resulting  $^{129}\text{I}/^{131}\text{I}$  atomic ratios are likely to be highly overestimated.

We, therefore, prefer the estimates based on the soil samples from Belarus. They are not affected by theoretical considerations. Belarus was the first country to receive the fresh fallout and together with other Eastern European states is less affected by pre-Chernobyl fallout. If the pre-Chernobyl fallout is small then the smallest  $^{129}\text{I}/^{131}\text{I}$  isotopic ratio should have the highest credibility. The mean isotopic ratio calculated from the work of Straume et al. (1986) and of Mironov et al. (2002) is  $^{129}\text{I}/^{131}\text{I} = 13.6 \pm 2.8$ , which we assume here as the best estimate.

On the basis of the total  $^{129}\text{I}$  inventories corrected for the pre-Chernobyl fallout of  $44 \text{ mBq m}^{-2}$ , the total  $^{131}\text{I}$  inventories were calculated using an initial isotopic ratio  $^{129}\text{I}/^{131}\text{I}$  of the Chernobyl fallout of  $13.6 \pm 2.8$  equivalent to an activity ratio  $A_{131}/A_{129} = (5.3 \pm 0.3) \cdot 10^7$ . The expectation values of the total  $^{131}\text{I}$  inventories for the villages investigated in this work are given in table 4.

#### **4.5 The aggregated dose coefficients for $^{131}\text{I}$ exposure**

The retrospective dosimetry of  $^{131}\text{I}$  exposures via  $^{129}\text{I}$  needs aggregated dose coefficients for  $^{131}\text{I}$  in units of Sv per Bq  $\text{m}^{-2}$  to convert the total  $^{131}\text{I}$  inventory (derived from total  $^{129}\text{I}$  inventories) into estimates of  $^{131}\text{I}$  thyroid equivalent doses and consequently depends on the reliability of radio-ecological modeling.

Robl et al. (1997) used the ECOSYS model (Müller and Pröhl, 1993) to calculate the inhaled and ingested  $^{131}\text{I}$  activities depending on the particular weather conditions during fallout. These authors emphasized that the main uncertainty in the model calculations depends on whether the fallout occurred via wet or dry deposition. The fact that the fallout in Northern Ukraine was mainly due to dry deposition (Talerko, 2004) makes this a minor problem. But, unfortunately, aggregated dose coefficients cannot be derived from the work of Robl et al. (1997).

Up to now, Pietrzak-Flis et al. (2003) were the only authors reporting actual  $^{131}\text{I}$  dose estimates on the basis of a retrospective dosimetry using  $^{129}\text{I}$ . They used the CLRP model (Krajewski, 1996) in the form of a computer code by Krajewski (1999) to derive the required dose coefficients. From the data given in their publication aggregated dose coefficients can be derived for 3 age groups. According to their model calculations, the dose coefficients for the committed thyroid equivalent doses are  $(2.5 \pm 0.1) \times 10^{-7} \text{ Sv Bq}^{-1} \text{ m}^2$  for 5-year-old children,  $(1.7 \pm 0.1) \times 10^{-7} \text{ Sv Bq}^{-1} \text{ m}^2$  for 10-year-old children, and  $(5.9 \pm 0.3) \times 10^{-8} \text{ Sv Bq}^{-1} \text{ m}^2$  for adults. We shall use them to give first order estimates of the thyroid doses in the Korosten and Narodici districts based on our integral  $^{129}\text{I}$  deposition densities.

It is to note, that these dose coefficients are much lower than those calculated by the German “Allgemeine Berechnungsgrundlage” (BMI, 1979) and the “Störfall-Berechnungsgrundlage” (BMI, 1977):  $1.3 \cdot 10^{-6} \text{ Sv Bq}^{-1} \text{ m}^2$  for a one-year-old child and  $0.22 \cdot 10^{-6} \text{ Sv Bq}^{-1} \text{ m}^2$  for an adult. These latter models are used in Germany for planning purposes and are known to give highly conservative results. The differences in the aggregated dose coefficients are, however, mainly due to differing model assumptions and only to a minor degree due to the differences between the differences in older (BMU, 1989) or newer ICRP dose coefficients which are part of the European Basic Safety Standards (CEC, 1996).

For more detailed dose estimates, fallout calculations taking into account the actual weather conditions in April and May 1986 have to be combined with realistic radioecological model calculations. Such calculations have not yet been performed. Calculations of the actual fallout fields were recently performed by Talerko (2004), but a radioecological modeling of the transfer from fallout to man considering the actual ecological conditions at the time of the accident is still missing. It is also beyond the scope of this publication.

#### 4.6 Exposure due to $^{131}\text{I}$ in villages of the Korosten and Narodici districts

In table 5, we present the results for the thyroid exposures based on the aggregated dose coefficients deduced from the work of Pietrzak-Flis et al. (2003). We did not apply a time correction for the decay of  $^{131}\text{I}$  from the reactor to the districts of Korosten and Narodici. There is official evidence that the radioactive cloud arrived already on April 27, 1986 at Narodici and that the ambient dose rate fell quickly after April 28, 1986.

According to our calculations, thyroid exposures due to  $^{131}\text{I}$  of the order of 10 Sv have to be assumed for the children living in 1986 close to Narodici where no countermeasures were undertaken at that time and where the population was not even warned to stay at home or to avoid drinking milk or eating fresh vegetables and salad. For adults the thyroid exposure was about 3 Sv. Also in the vicinity of Korosten thyroid doses of more than 1 Sv for children and of nearly 0.4 Sv for adults must be assumed.

All these dose estimates are in the range of thyroid doses (between 0.3 Sv und 40 Sv) derived from measurements of children aged up to 7 from Gomel in Belarus (NEA, 1995). The estimates are also compatible with the thyroid exposures of children from Belarus reported by UNSCEAR (2000) and are well within the range of individual thyroid dose measurements in Northern Ukraine (Goulko et al., 1998). Typically, the latter individual doses show in each age group a high variability with ranges of an order of magnitude or more. These ranges are well in line with the observed differences in  $^{129}\text{I}$  deposition densities with geometric standard deviations of the order of two. This variability calls for a probabilistic approach when calculating more refined dose coefficients for the retrospective dosimetry of  $^{131}\text{I}$  exposure using  $^{129}\text{I}$ . Presently, our data in table 5 just estimate the expectation value of the distribution of the individual thyroid doses.

Up to now,  $^{137}\text{Cs}$  was used more often to reconstruct the thyroid exposure to  $^{131}\text{I}$  than  $^{129}\text{I}$ . Kruk et al. (2004) summarized the respective earlier work and performed new model calculations for Belarus. They considered the actual weather and ecological conditions at the time of the accident and - as far as available - information about the fallout ratio of  $^{131}\text{I}$  and  $^{137}\text{Cs}$ . We used their results for aggregated dose coefficients for villages in Northern Ukraine where more than 10 measurements of the  $^{131}\text{I}$  and  $^{137}\text{Cs}$  fallout were available ( $3.6 \times 10^{-6}$  Sv per Bq  $\text{m}^{-2}$  for 5-year-old children,  $2.3 \times 10^{-6}$  Sv per Bq  $\text{m}^{-2}$  for 10-year-old children, and  $1.5 \times 10^{-6}$  Sv per Bq  $\text{m}^{-2}$  for adults) to convert  $^{137}\text{Cs}$  deposition densities to thyroid doses. The thyroid doses calculated in this way agree with our estimates based on  $^{129}\text{I}$  within a factor of two for adults and within 30 % for 5-years-old children. However, there remains the *caveat* that this agreement crucially depends on the knowledge of the ratio of the actual  $^{131}\text{I}$  and  $^{137}\text{Cs}$  fallout.

Thus, a retrospective assessment of the  $^{131}\text{I}$  exposure via  $^{129}\text{I}$  in soils appears feasible though the uncertainties of the derived doses are surely large. Surely direct measurements of the thyroids after the accident have to be preferred if available. For areas, however, where such measurements were lacking, retrospective dosimetry via  $^{129}\text{I}$  will be a valuable tool to assess the consequences of the accident.

## 5 Conclusions

After a nuclear accident, such as the Chernobyl catastrophe, the long-lived  $^{129}\text{I}$  in soils can be used to determine in retrospect the exposure of the thyroid to  $^{131}\text{I}$ . In the case of the Chernobyl accident, this is, however, only feasible in areas with high  $^{129}\text{I}$  inventories which statistically significantly exceed the pre-accident depositions from nuclear atmospheric weapons tests and from the emissions from European reprocessing plants. There remain large uncertainties for the thyroid dose estimates via  $^{129}\text{I}$  which are connected with still open questions about the migration of iodine isotopes in soils,

about the initial  $^{131}\text{I}/^{129}\text{I}$  activity ratio, and about the aggregated  $^{131}\text{I}$  dose coefficients. In spite of these open questions, which deserve further investigations, retrospective dosimetry via  $^{129}\text{I}$  can yield valuable information on the radiation exposure due to  $^{131}\text{I}$  in areas where no direct measurements of thyroids exist, in particular, since  $^{137}\text{Cs}$  cannot be used for this purpose.

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## Table and Figure Captions

Table 1: Results of RNAA for  $^{127}\text{I}$  and  $^{129}\text{I}$  in 51 samples of top soils ( $d \leq 10$  cm) from various affected and unaffected locations in Ukraine, Belarus, and Russia sampled between September 1991 and October 1992.

Table 2: Total inventories and their associated standard uncertainties of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  in Moscow and in three different areas of Northern Ukraine.  $^{137}\text{Cs}$  data are decay-corrected for April 26, 1986. The entries under “sample type” indicate whether the inventories were determined from measured depth profiles or from average samples.

Table 3: Geometric means and standard deviations of total  $^{129}\text{I}$  inventories as derived from the analysis of soil profiles from differently contaminated areas.

Table 4:  $^{129}\text{I}/^{131}\text{I}$  atomic ratios measured in environmental samples after the Chernobyl accident.

Table 5: Expectation values  $E(D_{\text{corr}}(\text{I-129}))$  according to equation 6 of total inventories of  $^{129}\text{I}$  measured in zone II and III corrected for global weapons fallout, inferred expectation value of total  $^{131}\text{I}$  inventories  $E(D(\text{I-131}))$  and resulting  $^{131}\text{I}$  radiation exposures of thyroid glands  $H_{\text{thyroid}}$  using the aggregated dose factors  $DC_{131}$  according to Pietrzak-Fils et al. (2003).

Fig. 1: Map of Northern Ukraine with contamination zones II ( $^{137}\text{Cs}$  deposition density  $> 555$   $\text{kBq m}^{-2}$ ) and III ( $^{137}\text{Cs}$  deposition density  $185 - 555$   $\text{kBq m}^{-2}$ ) and the sampling sites. The circles indicate 30 km and 60 km distances from the NPP, respectively.

- Fig. 2: Concentrations in  $\text{mg kg}^{-1}$  (top) and inventories within individual depth increments in  $\text{mg m}^{-2}$  (bottom) of  $^{127}\text{I}$  in soil profiles from Moscow, Nemirovka, and Nosdrischtsche.
- Fig. 3: Depth profiles of  $^{129}\text{I}$  specific activities in soils from Moscow (Moscow VI: diamonds; Moscow VII: squares) and Northern Ukraine: zone IV (Lewkow 1: triangles; Lewkow 2: diamonds), zone III (Nemirovka 2: open triangles; Woronewo 4: squares; Kupetsch-Tschernjanka 1: diamonds), and zone II (Nosdrischtsche II: diamonds; Christinovka meadow: open circles; Christinovka river shore: squares).
- Fig. 4:  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios in the soil profiles Moscow VI (diamonds), and Moscow VII (squares), Nemirovka II (triangles) and Nosdrischtsche 2 (circles).
- Fig. 5: Inventories within individual depth increments of  $^{129}\text{I}$  (full symbols) and  $^{137}\text{Cs}$  (open symbols) in soils from (top panel) Moscow VI (diamonds), Moscow VII (squares) and from Levkov (circles) and Barashevka (triangles) in Ukraine; (middle panel) contamination zones III (Tschigiri: full circles; Kupetsch-Tschernjanka: full squares; Woronewo: full diamonds; Nemirovka: full triangles) and (bottom panel) contamination zone II (Christinovka river shore: squares; Christinovka meadow: triangles).
- Fig. 6: Total  $^{129}\text{I}$  and  $^{137}\text{Cs}$  inventories measured from this work (full symbols) in depth profiles from Moscow/Russia (diamonds), Zhitomir region/Ukraine (circles), Korosten region/Ukraine (squares) and Narodici region/Ukraine (triangles). For comparison, the results of the investigations by Straume et al. (1996) (open triangles) for Belarus and by Mironov et al. (2002) for Belarus (open diamonds) and the logarithmic mean and

standard deviation (cross) obtained for seven soil profiles from Lower Saxony/Germany by Ernst et al. (2003) are given.

Table 1: Results of RNAA for  $^{127}\text{I}$  and  $^{129}\text{I}$  in 51 samples of top soils ( $d \leq 10$  cm) from various affected and unaffected locations in Ukraine, Belarus, and Russia sampled between September 1991 and October 1992.

	I-127 in mg/kg	I-129 in mBq/kg	129/127 in $10^{-10}$
Geometric mean and standard deviation	$1.9 \times 2.2^{\pm 1}$	$3.0 \times 3.3^{\pm 1}$	$2700 \times 4.7^{\pm 1}$
Minimum	0.29	0.05	45
Maximum	8.6	111	74000

Table 2: Total inventories and their associated standard uncertainties of  $^{129}\text{I}$  and  $^{137}\text{Cs}$  in Moscow and in three different areas of Northern Ukraine.  $^{137}\text{Cs}$  data are decay-corrected for April 26, 1986. The entries under “sample type” indicate whether the inventories were determined from measured depth profiles or from average samples.

location	sample type	$^{129}\text{I}$ analytical method	$^{129}\text{I}$		$^{137}\text{Cs}$	
			D in mBq/m <sup>2</sup>	u(D)	D in kBq/m <sup>2</sup>	u(D)
Moscow VI 1996	profile	AMS	64	5	4.57	0.37
Moscow VII 1996	profile	AMS	37	3	3.63	0.29
<b>not contaminated</b>						
Barashevka 1	profile	RNAA/AMS	32.1	9.1	3.14	0.30
Barashevka 2	average	AMS	25.4	2.8	5.88	0.47
Barashevka 3	average	AMS	34.2	3.8	4.28	0.35
Barashevka 4	average	AMS	29.8	5.1	3.69	0.33
Oserjanka 1	average	AMS	44.2	4.6	6.36	0.50
Oserjanka 2	average	AMS	17.7	1.9	3.57	0.31
Oserjanka 3	average	AMS	34.4	3.6	6.79	0.40
Davidovka 1	average	AMS	28.6	3.1	4.64	0.41
Davidovka 2	average	AMS	47.4	4.9	4.32	0.40
Levkov 1	profile	RNAA/AMS	81.6	8.2	5.27	0.46
Levkov 2	profile	RNAA/AMS	101	11	6.22	0.47
<b>medium contaminated (zone III)</b>						
Nemirovka 2	profile	RNAA/AMS	186	14	492	44
Nemirovka 3	average	AMS	237	25	454	50
Nemirovka 4	average	AMS	74.4	7.8	394	51
Woronewo 1	average	AMS	195	32	527	54
Woronewo 2	average	AMS	116	12	375	41
Woronewo 3	average	AMS	172	31	356	84
Woronewo 4	profile	RNAA/AMS	218	24	851	110
Woronewo 5	average	AMS	175	19	612	60
Woronewo 6	average	AMS	181	19	598	55
Woronewo 7	average	AMS	85.7	9.2	296	19
Kupetsch-Tschernjanka 1	profile	RNAA/AMS	197.	19	332	31
Kupetsch-Tschernjanka 2	average	AMS	100	11	277	21
Tschigiri 1	profile	RNAA/AMS	300	29	320	37
Tschigiri 2	average	AMS	78.8	8.1	338	34
Tschigiri 3	profile	RNAA/AMS	138	16	397	41
Tschigiri Zwintor 1	average	AMS	88.2	8.0	356	43
Tschigiri Zwintor 2	average	AMS	168	15	237	23
Tschigiri Zwintor 3	average	AMS	97.0	8.8	528	52
Tschigiri Rjet. Most 2	average	AMS	81.5	8.4	350	38
Tschigiri Rjet. Most 3	average	AMS	139	14	556	51
Tschigiri Ferma 1	average	AMS	105	11	381	36
Tschigiri Ferma 2	average	AMS	72.7	7.5	360	39
Tschigiri Chmilnik 1	average	AMS	111	11	532	52
Tschigiri Chmilnik 2	average	AMS	95.6	9.9	442	39
<b>highly contaminated (zone II)</b>						
Nosdrischtsche 2	profile	AMS	1390	100	5229	327
Nostrischtsche 3	average	AMS	653	71	2587	1382
Novo Scharno 2	average	AMS	974	97	3318	237
Novo Scharno 3	average	AMS	811	80	3479	276
Novo Scharno 4	average	AMS	1212	120	3916	457
Christinovka river shore	profile	RNAA/AMS	1098	122	4218	436
Christinovka meadow	profile	RNAA/AMS	388	44	738	50



Table 3: Geometric means and geometric standard deviations of total  $^{129}\text{I}$  inventories as derived from the analysis of soil profiles from differently contaminated areas.

location and year of sampling	number of profiles	depth of profiles in cm	total $^{129}\text{I}$ inventory in $\text{mBq m}^{-2}$	reference
pre-nuclear				
Lutovinovo, Russia, 1939	1	35	$0.084 \pm 0.017$	Szidat et al. (2000b)
not contaminated by fall-out from Chernobyl				
Moscow, Russia, 1996	2	40	$49 \times 1.5^{\pm 1}$	this
Zhitomir, Ukraine, 1997	12	40	$38 \times 1.7^{\pm 1}$	this
Lower Saxony, D, 1999	7	250	$168 \times 1.5^{\pm 1}$	Ernst et al. (2003)
contamination zones in Ukraine				
mean, zone III, 1995	24	40	$130 \times 1.5^{\pm 1}$	this
Nemirovka, 1995	3	40	$94 \times 2.1^{\pm 1}$	this
Woronowo, 1995	7	40	$109 \times 1.7^{\pm 1}$	this
Kupetsch-Tschernjanka, 1995	2	40	$93 \times 2.0^{\pm 1}$	this
Tschigiri, 1995	12	40	$67 \times 1.9^{\pm 1}$	this
mean, zone II, 1995	7	40	$848 \times 1.5^{\pm 1}$	this
Nosdrischtsche, 1995	2	40	$905 \times 1.8^{\pm 1}$	this
Novo Scharno, 1995	3	40	$941 \times 1.2^{\pm 1}$	this
Christinovka, 1995	2	40	$602 \times 2.2^{\pm 1}$	this

Table 4:  $^{129}\text{I}/^{131}\text{I}$  atomic ratios measured in environmental samples after the Chernobyl accident. The  $^{131}\text{I}$  data are given for April 26, 1986.

material	origin	sampling date	$^{129}\text{I}/^{131}\text{I}$ atomic ratio <sup>a)</sup>	reference
sewage sludge	Berlin/D	May 1986	$34.3 \pm 2.8$	this work
bovine and deer thyroid glands	Ulm/D, Bad Hall/A	May - June 1986	$27 \pm 10$	VanMiddlesworth and Handl (1997)
rain	Munich/D	end of April 1986	$19 \pm 5$	Kutschera et al. (1988)
rain	Israel	5. May 1986	$15 \pm 3$	Paul et al. (1987)
soil	Belarus	May - June 1986	$15 \pm 5$	Mironov et al. (2002)
soil	Belarus	May 1986	$12 \pm 3$	Straume et al. (1996)
theoretical estimate	-	-	11 - 15	Ermilov (1993)
theoretical estimate	-	-	22.8	Kirchner and Noack (1988)

a)  $^{131}\text{I}$  concentrations are decay-corrected for April 26, 1986.

Table 5: Expectation values  $E(D_{\text{corr}}(\text{I-129}))$  according to equation 6 of total inventories of  $^{129}\text{I}$  measured in zone II and III corrected for global weapons fallout, inferred expectation value of total  $^{131}\text{I}$  inventories  $E(D(\text{I-131}))$  and resulting  $^{131}\text{I}$  radiation exposures of thyroid glands  $H_{\text{thyroid}}$  using the aggregated dose factors  $DC_{131}$  according to Pietrzak-Fils et al. (2003).

	$E(D_{\text{corr}}(\text{I-129}))$ in $\text{Bq m}^{-2}$	$E(D(\text{I-131}))$ in $10^6 \text{Bq m}^{-2}$	committed equivalent thyroid dose $H_{\text{thyroid}}$ in Sv		
			5-year-old	10-years-old	adult
medium contaminated area (zone III) close to Korosten					
Nemirovka	0.122	6.41	1.6	1.1	0.38
Woronewo	0.124	6.50	1.6	1.1	0.38
Kupetsch	0.119	6.25	1.5	1.0	0.37
Tschigiri	0.078	4.10	1.0	0.7	0.24
highly contaminated area (zone II) close to Naroditschi					
Nosdrischtsche	1.060	55.7	14	9.3	3.3
Novo Scharno	0.962	50.5	13	8.4	3.0
Christinovka	0.824	43.3	11	7.2	2.6

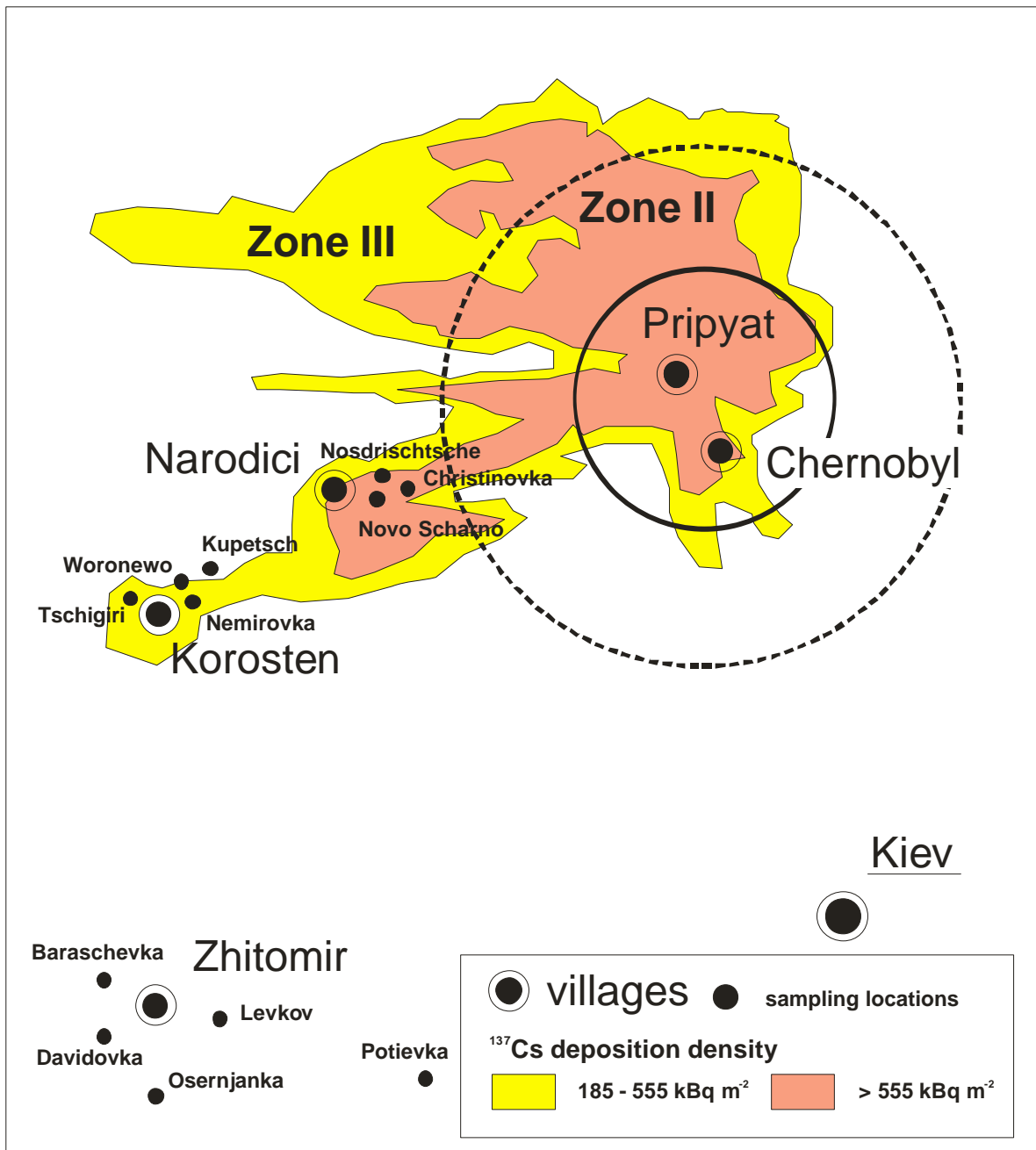


Fig. 1: Map of Northern Ukraine with contamination zones II ( $^{137}\text{Cs}$  deposition density  $> 555 \text{ kBq m}^{-2}$ ) and III ( $^{137}\text{Cs}$  deposition density  $185 - 555 \text{ kBq m}^{-2}$ ) and the sampling sites. The circles indicate 30 km and 60 km distances from the NPP, respectively.

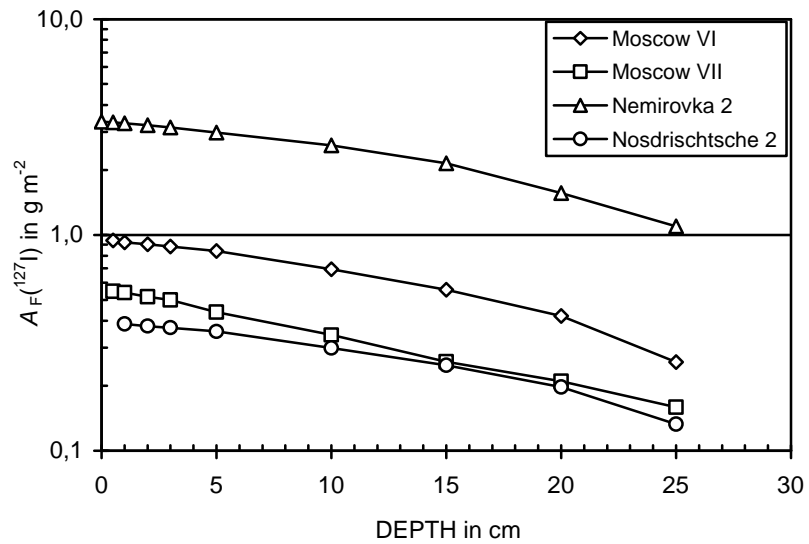
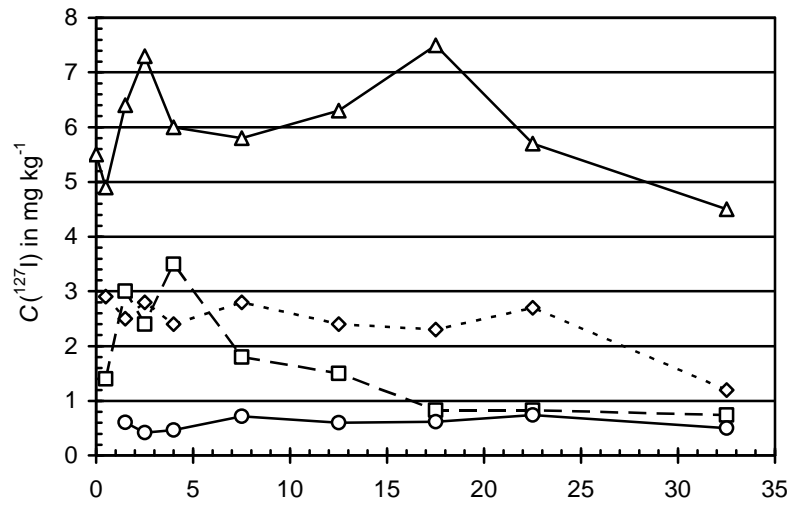


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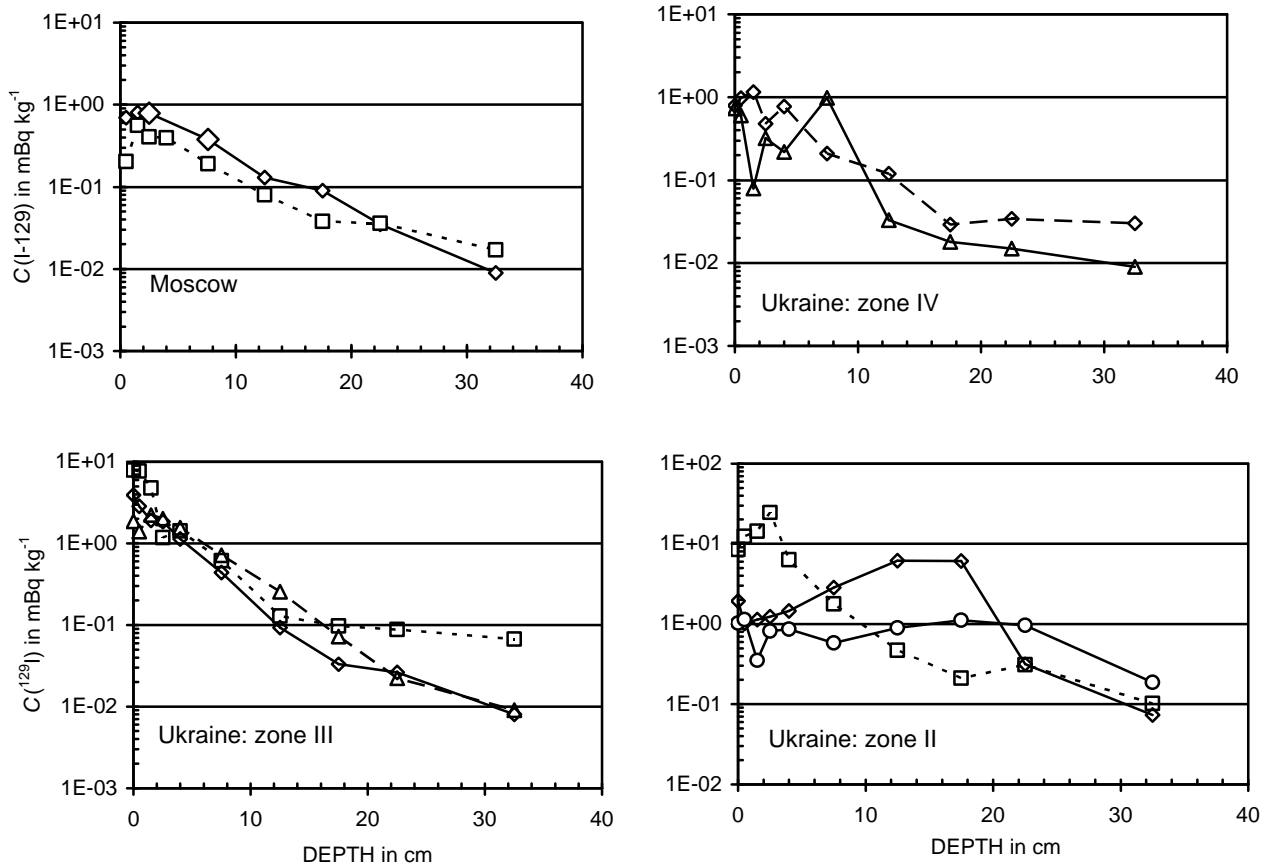


Fig. 3: Depth profiles of  $^{129}\text{I}$  specific activities in soils from Moscow (Moscow VI: diamonds; Moscow VII: squares) and Northern Ukraine: zone IV (Levkov 1: triangles; Levkov 2: diamonds), zone III (Nemirovka 2: open triangles; Woronewo 4: squares; Kupetsch-Tschernjanka 1: diamonds), and zone II (Nosdrischtsche II: diamonds; Christinovka meadow: open circles; Christinovka river shore: squares).

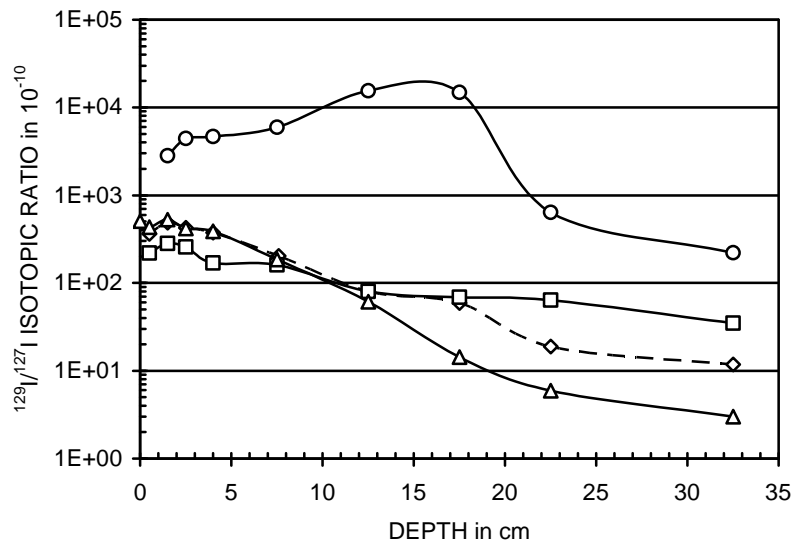


Fig. 4:  $^{129}\text{I}/^{127}\text{I}$  isotopic ratios in the soil profiles Moscow VI (diamonds), and Moscow VII (squares), Nemirovka II (triangles) and Nosdrischtsche 2 (circles).

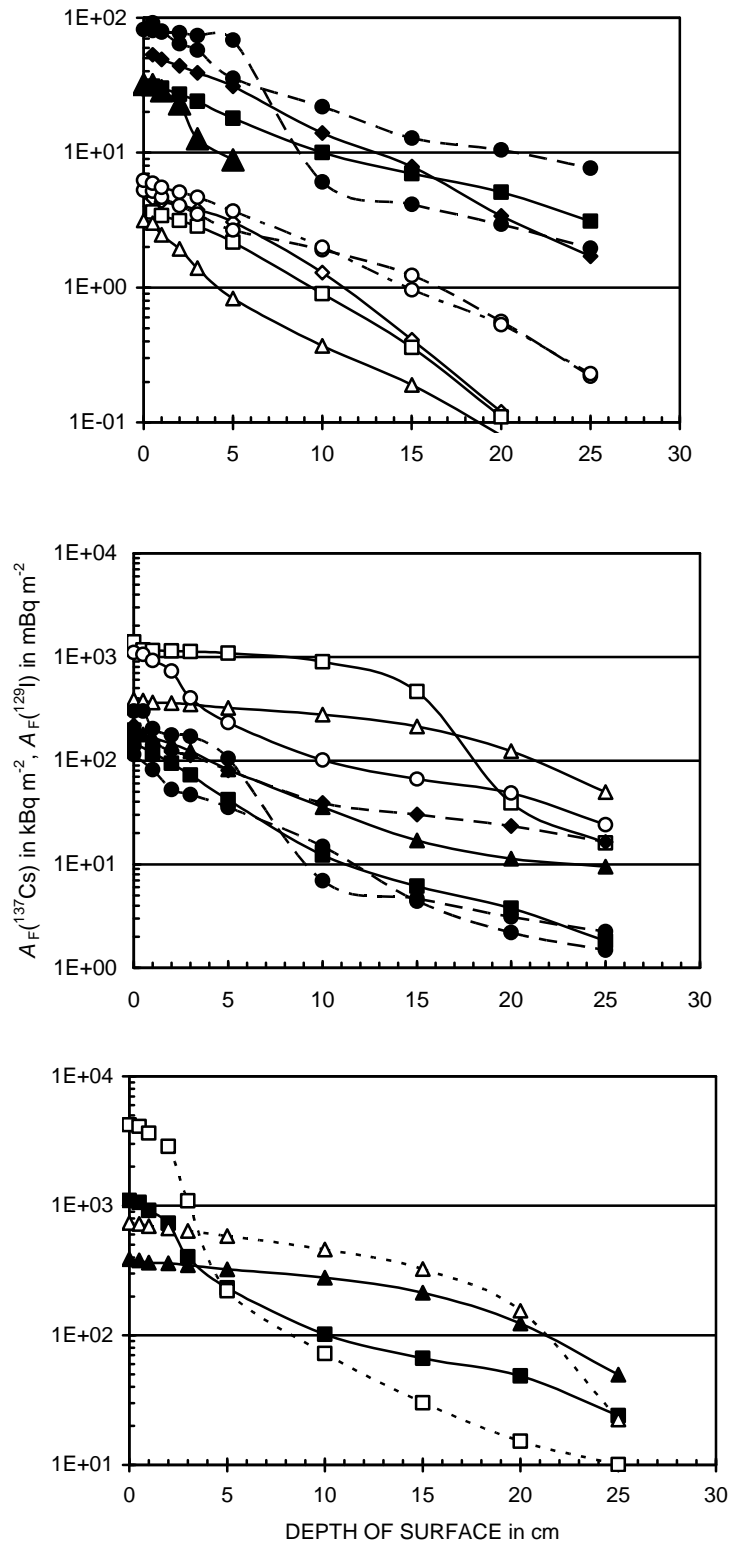


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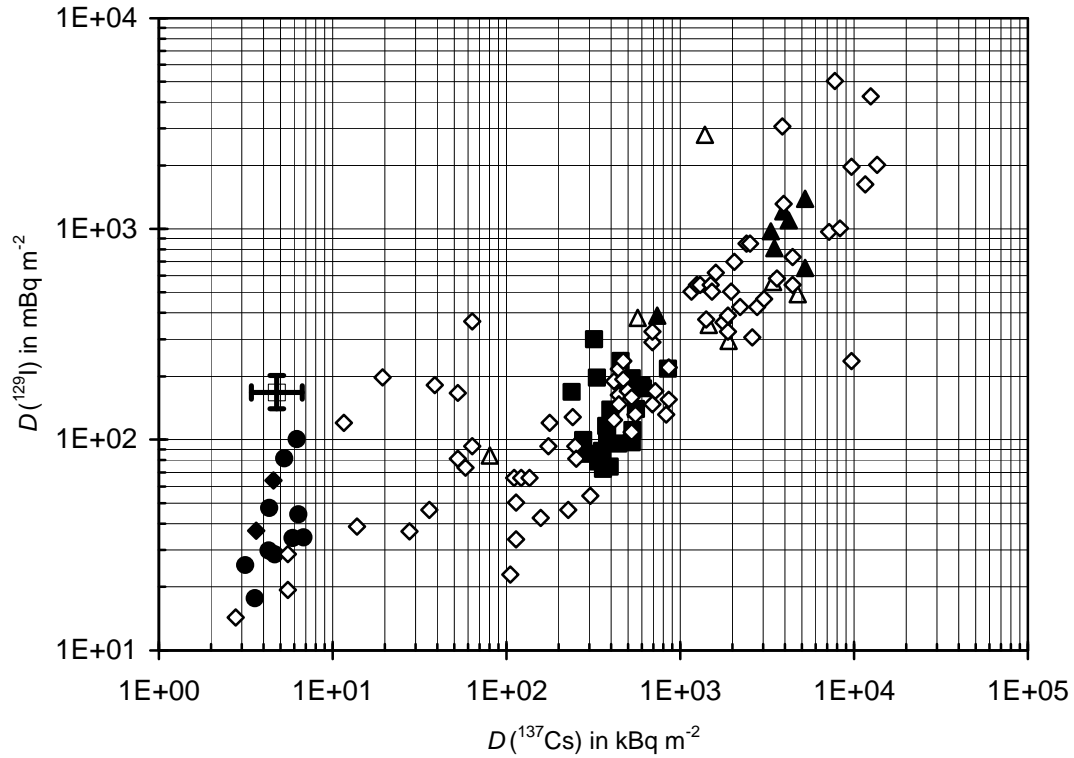


Fig. 6: Total  $^{129}\text{I}$  and  $^{137}\text{Cs}$  inventories from this work (full symbols) measured in depth profiles from Moscow/Russia (diamonds), Zhitomir region/Ukraine (circles), Korosten region/Ukraine (squares) and Narodici region/Ukraine (triangles). For comparison, the results of the investigations by Straume et al. (1996) (open triangles) for Belarus and by Mironov et al. (2002) for Belarus (open diamonds) and the geometric means and geometric standard deviations (cross) obtained for seven soil profiles from Lower Saxony/Germany by Ernst et al. (2003) are given.